

An intercomparison study involving eight long-range transport models for sulfur deposition in East Asia has been initiated. The participating models included Eulerian and Lagrangian frameworks, with a wide variety of vertical 31 resolutions and numerical approaches. Results from this study, in which models used common data sets for emissions, meteorology, and dry, wet and chemical conversion rates, are reported and discussed. Model results for sulfur dioxide 33 and sulfate concentrations, wet deposition amounts, for the period January and May 1993, are compared with observed quantities at 18 surface sites in East Asia. At many sites the ensemble of models is found to have high skill in predicting 35 observed quantities. At other sites all models show poor predictive capabilities. Source-receptor relationships estimated by the models are also compared. The models show a high degree of consistency in identifying the main source-receptor 37 relationships, as well as in the relative contributions of wet/dry pathways for removal. But at some locations estimated deposition amounts can vary by a factor or 5. The influence of model structure and parameters on model performance is 39 discussed. The main factors determining the deposition fields are the emissions and underlying meteorological fields. Model structure in terms of vertical resolution is found to be more important than the parameterizations used for 41 chemical conversion and removal, as these processes are highly coupled and often work in compensating directions. © 2001 Published by Elsevier Science Ltd. 43

- 45 Keywords: Model intercomparision; Sulfur deposition; Long-range transport; Source-receptor relationships
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49 1. Introduction

51 Fueled by high population growth and expanding economies, energy consumption in Asia is estimated to 53

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(G.R. Carmichael).

grow to 30% of the worlds total by 2015 (Shah et al.,
2000). Since fossil fuels will continue to provide much of
this energy, emissions of greenhouse gases and air
pollutants such as sulfur and nitrogen oxides and
particulates are projected to dramatically increase. The
impacts of Asia's growth in emissions will have wide-
ranging consequences (Streets et al., 1999). Acid
precipitation is an illustrative example. China's National57

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- 1 Environmental Protection Agency (NEPA) recently released a report indicating that economic losses due
- 3 to acid rain damage to forests and farmland are five times higher than initially assessed in 1996, and
- are now estimated at \$13.25 billion annually (Shah et al., 2000). The long-range transport and fate of pollutants in
- 7 Asia is an area of increasing scientific interest and political concern, as countries receive growing amounts9 of pollutants from neighboring and even distant
- countries.
 Long-range transport models play a critical role in science and policy analysis. These computer-based
- 13 models calculate the distribution of trace gases in the troposphere and the quantities of acids deposited on the surface from specified emissions distributions and
- meteorological scenarios. The major features of such models consist of: (1) a transport component (or
- module) to describe the wind speed and direction, the eddy diffusivity and mixing layer height, the tempera-
- ture, the water vapor, cloud water content, and the radiation intensity of each location as a function of time;
- (2) a chemical kinetic mechanism to describe the rates ofatmospheric reactions, including homogeneous gas-phase, heterogeneous, and liquid-phase reactions; and
- (3) removal modules to describe the dry deposition of material, and the in-cloud and below-cloud removal
- 27 processes. Each process incorporated into a model is itself a very complex and incompletely understood
- 29 phenomenon. Therefore, in formulating such models it is necessary to incorporate the processes into the
- model framework by utilizing chemical, dynamic, and thermodynamic parameterizations. Furthermore,
 even processes that are well understood may require
- parameterization to maintain some balance of the details among the different processes that are treated in the model and the availability and quality of the supporting
- data needed to implement and evaluate the model components. In addition there are various approaches
 to modeling long-range transport (i.e., Lagrangian and
- Eulerian frameworks). Each model has its own strengths and weaknesses.
- The potential for transboundary movement of sulfur
 and other pollutants means that domestic energy needs
 will become of increasing environmental concern (Shah
 et al., 2000; Guttikunda et al., 2001). Long-range
 transport models will be used to provide information
 on the transport and fate of emissions from various
 locations in East Asia. Source-receptor relationships in
 East Asia have been investigated by Arndt et al. (1998),
- Huang et al. (1995) and Ichikawa and Fujita (1995), 51 Calori et al. (2001) and the results are very different.
- The calculated contribution of Chinese sources to Japan's deposition in these studies present markedly
- different estimates of the role that long-range transport 55 plays in Japan's over-all deposition. Huang et al.
- (1995) estimated that China accounts for only 3.5% of

Japan's total sulfur deposition. They found that 57 over 93% of the sulfur deposited within Japan was from either Japanese anthropogenic or volcanic sources. 59 In contrast, Ichikawa and Fujita (1995) estimated China to be a major source of wet sulfate deposition 61 in Japan, accounting for one-half of the anthropogenic deposition. These variations are due in part 63 to differences in removal rates and chemical conversion rates in different models. Low removal rates 65 result in greater transport away from source locations and thus higher transboundary pollution. There is a 67 great need to conduct model intercomparision studies to better understand how to model long-range transport in 69 East Asia.

Previous work on intercomparing acid deposition71models applied to Japan has been conducted. Theseinclude a joint study comparing CRIEPI (Central73Research Institute of Electric Power Industry;Ichikawa and Fujita, 1995) and ATMOS (Arndt and75Carmichael, 1995) trajectory models with Eulerian77results (i.e., the STEM model, Carmichael et al., 1991).77These results have been reported by Phadnis and79

A Workshop on the Transport of Air Pollutants in Asia was held at International Institute for 81 Applied System Analysis (IIASA) in Laxenburg, Austria during 27-29 July 1998 (IIASA, 1998), and issues 83 related to the modeling of long-range transport in Asia were discussed. The attendees recognized that in order 85 to help improve the use of models in science and policy analyses in Asia it is necessary to have a 87 better understanding of model performance and uncertainties. The group initiated a model intercomparison 89 exercise to be performed as part of joint collaborative studies involving the CRIEPI/IIASA and the Regional 91 Air Pollution INformation and Simulation (RAINS)-Asia projects. The results from this study are presented 93 in this paper.

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2. Intercomparison framework

2.1. Task subdivision and conditions

The study domain (Fig. 1) 95-150°E and 5-50°N, 101 includes Southeast Asia, Philippines, Taiwan, Japan, North and South Korea, most of China, and parts of 103 Russia and Mongolia. Two month-long periods were selected (January and May 1993), to reflect long-range 105 transport conditions under two distinct seasons. During these periods measurements of sulfur concentrations and 107 deposition were made throughout the study region using identical sampling and analysis protocols. A common 109 meteorological data set was also prepared, using Regional Atmospheric Modeling System (RAMS) mod-111 el, and provided to all participants.

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The SO_2 emissions, based on the RAINS-Asia analyses, were supplied to participants through a web site (CGRER, 1999). They included emissions from area

- 3 site (CGRER, 1999). They included emissions from area sources and large point sources, emissions from shipping
- activities and from volcanoes. Release heights were also specified: area emissions were to be injected in the first
 model layer; while large point sources and volcanoes
- were assigned effective release heights of 300 and 9 1500 m, respectively. Emissions were considered as 95% SO₂ and 5% SO₄.
- 11 Eulerian models were run with zero initial and boundary conditions for sulfur and were initialized for
- 48 h. Prescribed deposition velocities were also used, as specified in Table 1 and the related land/sea mask was also made available.



Fig. 1. Modeling domain with countries/regions sets for source-receptor calculations, measurement site locations (●) and receptor locations (×): (1) Kangwha; (2) Komae; (3)
Taichung; (4) Nanjing; (5) Jinan; (6) Beijing; (7) Nangoku; (8) Otobe; (9) Amami; (10) Kashima; (11) Yangyang; (12) Tokoro;
(13) Hachijo; (14) Oki; (15) Tsushima; (16) Fukue; (17) Miyako; (18) Taitong.

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45 Table 1

	Prescribed	deposition	velocities
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Species	Conditions	$v_d \ (\mathrm{cm} \ \mathrm{s}^{-1})$
SO ₂	Land (April–September)	0.25
	Land (otherwise)	0.125
	Sea	0.32
SO_4	Land	0.2
	Sea	0.1

^a These parameters were selected as a result of a review of
 current literatures. Details can be found in Guttikunda et al.
 (2001) and Calori et al. (2001).

2.1.1. Intercomparison tasks

The intercomparison activity consisted of four tasks. All participants were asked to do Task A, and as many of the other tasks as possible.

2.1.1.1. Task A-the blind test. Modelers were requested to use their models and the above information 63 on emissions, meteorology, and dry deposition velocities, and calculate surface SO2 and sulfate concentra-65 tions and sulfur deposition. They were allowed to use their best estimate of model parameters associated with 67 chemical conversion and wet removal. Their results were reported for specific measurement site locations (Fig. 1) 69 in terms of daily average concentrations and depositions. These data were compared to measurement values 71 centrally by the study coordinators. Concentrations as well as wet and dry deposition fields at given times were 73 also reported (see below for details).

Every model result was blind (i.e., only the organizers75and the owner of the model knew their code number)71and all results were reported using the code numbers. In77addition each modeler submitted detailed information79regarding model parameters.79

2.1.1.2. Task B—the fixed parameter test.Extension of81Task A, but modelers were given specific parametervalues (Table 2) to use for SO_2 to sulfate conversion and83wet removal of SO_2 and sulfate.Submitted results were85the same as before.85

2.1.1.3. Task C—the source-receptor test. Modelers
submitted source-receptor information for a few prescribed receptor locations, selected from the measurement sites: Komae, Oki, Fukue, Yangyang, Beijing, Nanjing, Taichung. Contributions from various sets of
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Table 2	95
Prescribed parameters for Task B ^a	
	(1)/

Parameter	Value	91
$SO_2 \rightarrow SO_4$ conversion rate (s ⁻¹)	$K_0 f(\lambda) + K_1 [1 - f(\lambda)] g(J)$	99
	where	101
	$K_0 = 1 imes 10^{-5}$ $K_1 = 4 imes 10^{-6}$	101
	$f(\lambda) = \cos[1.3\lambda\pi/180]$	102
	$g(J) = \sin[(J - 80)2\pi/365]$	105
SO ₂ wet removal rate	$2 \times 10^{-5} P$	
constant (s^{-1})		105
SO ₄ wet removal rate	$5 imes 10^{-5} P^{0.83}$	
constant (s ⁻¹)		107

^a These parameters were selected as a result of a review of current literatures. Details can be found in Guttikunda et al. (2001) and Calori et al. (2001).

 λ is latitude (°N), J is the Julian day and P is the 111 precipitation rate (mm h⁻¹).

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- countries/regions (Fig. 1) and volcanic sources were requested.
- 2.1.1.4. Task D—the tuning test. Modelers were given
 the observation data and given the chance to rerun their model (changing parameters as they saw fit) and submit
 new results.

9 2.1.2. Model results reporting procedures

Reporting activities followed five stages, according to 11 the task subdivisions. At each stage of the intercompar-

- ison participants were requested to submit data as well as relevant information to help in the interpretation of
- model results. The reporting procedures are outlined in Table 3. A code number was assigned to each registered participants, and was used when submitting model
- 17 results.

19 2.2. Meteorological data

The study region has dramatic variation of topography, land type, and mixtures of industrial/urban centers
and agricultural/rural regions. It also includes a large portion of ocean (more than 30% of the domain), where
there is a void of in situ meteorological data and so the use of numerical models is crucial to capture the
interactions between continental and marine air masses

determining pollutants transport, transformation and deposition in the region. A common meteorological data

set was prepared by Prof. Uno of Kyushu University 31 (Uno and Jang, 1999) using the RAMS model (Pielke

- et al., 1992). RAMS includes the Kuo-type cumulus parameterization to represent the sub-grid scale convective cumulus and the Kessler-type microphysics model, able to simulate mesoscale cloud and precipita-
- tion phenomena. The surface flux calculation module in RAMS (Louis, 1979) was improved based on the result
- of Uno et al. (1995). The employed computational grid 39 consists of 60×56 grid cells of 100 km each in horizontal, and 30 vertical levels following the topo-
- 41 graphy, with the top level set as 23 km. Model field initialization and four-dimensional data assimilation
- 43 (FDDA) were performed using a Newtonian relaxation ('nudging') based on the time-dependent lateral bound-
- 45 ary conditions provided based on the European Centre for Medium-Range Weather Forecasts (ECMWF) 2.5°
- 47 global analysis data. The initialization was performed 3 days before each monthly period and the nudging was
 49 prescribed every 12 h for 5 lateral grids.
- At the ground surface, topography and the land/sea 51 mask were obtained from the RAMS built-in data set, and the vegetation index was converted from Matthews
- 53 data. (Matthews, 1983) Sea surface temperatures (SST) were taken from (National Center for Environmental
- 55 Prediction (NCEP), US) database (Vazquez et al., 1998, $1^{\circ} \times 1^{\circ}$ resolution), while snow cover fraction was taken

from International Satellite Land Surface Climatology57Project (ISLSCP) Initiative I data set (Sellers et al., 1995;59weekly data in 1987). The stored meteorological variables (every 3 h) distributed to the modelers included59three-dimensional fields of wind components, air temperature, atmospheric pressure, water vapor mixing61ratio, cloud water and rainwater, and 2D fields of63surface temperature, friction velocity and temperature,65

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2.2.1. General meteorological conditions

Fig. 2 shows the main meteorological features of the two simulated months. Winter monsoon (WM) is the 69 weather type mostly characterizing the month of January, typically associated with high wind speed 71 determining transboundary pollutant transport eastward of mainland Asia. Meteorological fields calculated 73 by RAMS for January 1993 indicate the occurrence of transition weather types before the establishment of the 75 WM. A Japan Sea low type (JSL) is observed from 2 to 4 January, while the CRIEPI observation period (11-21 77 January) is characterized by a sequence of different weather types: a typical winter monsoon type (WM, 11 79 January), followed by a South Coast low type (SCL, 12-16 January), a WM (16-20 January) and then again a 81 SCL (23–26 January).

In May, the typical weather is characterized by the large scale traveling high pressure system moving very slowly eastward. However, in May 1993 such a system did not occur; rather only a small high speed traveling high pressure system was observed (i.e., 4–5, 16–17 May), with an intermediate west-east meso-front on 9– 10 May. 89

The comparison between calculated and observed precipitation at monitoring sites on a monthly basis is presented in Fig. 3.

2.3. Observational data

Observational data used in this study were those obtained at the 18 stations (shown in Fig. 1) comprising 97 the cooperative monitoring network for acid deposition in East Asia. The network was operated by the Institute 99 of Water Conservancy and Hydroelectric Power Research (IWHR) in mainland China, Kon-Kuk Univer-101 sity (KKU) in Korea, Taiwan Power Company in Taiwan, and the Central Research Institute of Electric 103 Power Industry in Japan, for 3 years from October 1990. A major purpose of the monitoring was to build a data 105 set of acid deposition for long-term, regional-scale evaluation in East Asia. The stations were carefully 107 selected considering factors like climatic precipitation amounts, geology, source location, and convenience of 109 maintenance, based on Robertson and Wilson (1985). See Fujita and Takahashi (1994) and Fujita et al. (2000) 111 for more details.

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1 Table 3

Outline of reporting procedures

Stage	Activities
Registration	All the groups participating to the intercomparison exercise were encouraged to register via e-mail, and to provide the following information:
	 participant information (name, organization, planned contributions) model structure and purpose: model name, main model focus, modelled quantities, model type and dimension
Task A The blind test	Data requested
	 monitoring sites data set: daily averaged surface concentrations of SO₂ and SO₄ and daily dry and wet sulfur (SO₂ + SO₄) depositions at the specified measurement site locations (see Fig. 1) daily averaged surface concentrations fields: 2D fields of daily averaged concentrations of SO₂ and
	 SO₄ at the surface on every 10th day (i.e., 10 January, 20 January, 30 January, 10 May, etc.) <i>daily averaged boundary layer concentrations fields</i>: 2D fields of daily averaged concentrations of SO₂ and SO₄, averaged from the ground to 1 km height on every 10th day (data set not required if
	 identical to the one of daily averaged surface concentrations) <i>hourly averaged surface concentrations fields</i>: 2D fields of hourly averaged concentrations of SO₂
	 IO days averaged surface concentrations fields: 2D fields of 10 days averaged concentrations of SO₂ and SO₄ at 00 UTC hour on every 10th day
	• 10 days accumulated dry and wet depositions fields: 2D fields of 10 days accumulated SO ₂ and sulfate dry and wet depositions (i.e. 1–10 January, 11–20 January)
	Accompanying information A description of model characteristics and modeling conditions:
	• <i>model domain</i> : horizontal and vertical size and location of the domain, horizontal and vertical and vertical and vertical size and goid system
	 <i>physical and chemical processes</i>: basic assumptions, numerical schemes and adopted parameters for each process: transport and diffusion, chemistry, dry deposition, wet removal
	• input data used:
	 copography emissions: species list, spatial resolution, time frequency, sources types (area/point) meteorology: source (e.g. observations vs. model), spatial resolution, time frequency, variables list
	 initial and boundary conditions: source (e.g. observations, other models), type (e.g. fixed, variable in time)
	• <i>emissions summary</i> (if an emission data set different from the one made available for the intercomparison, a table with total emissions by country)
	• <i>references</i> : all relevant references (preferably in open literature), pertaining to model formulation and previous applications
Task B The fixed parameter test	Data requested Same data sets as for Task A
	Accompanying information A detailed description of any parameterizations used in the model different from those prescribed
Task C The source-receptor test	Data requested
	<i>source-receptor relations</i> : contributions to total depositions from sets of countries/regions (Fig. 1) and volcanic sources at the prescribed receptor locations: Komae, Oki, Fukue, Yangyang, Beijing, Nanjing, Taichung
Task D The tuning test	Monitoring data have been provided to each participant after submission of model results for Tasks A, B and C
	Data requested Same data sets as for Task A
	Accompanying information adopted parameters for each of the following processes: chemical mechanism, dry deposition, wet removal

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Fig. 2. Upper panels: monthly mean modeled winds at 1000 m sigma level: (a) January; (b) May; lower panels: monthly modeled total precipitation (mm): (c) January; (d) May.

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Precipitation samples were collected with wet-only 35 samplers with an aperture area of $190 \,\mathrm{cm}^2$. The top of the sampler was at approximately 150 cm above the 37 ground level. The precipitation sensor of the sampler responded to rain drops with a diameter larger than 39 0.5 mm. Heaters were thermostatically controlled to prevent freezing and snow accumulation, when tem-41 peratures fell below 0°C. The amount of precipitation was measured using tipping-bucket rain gauges with a detection limit of 0.5 mm. Precipitation samples ob-43 tained at stations in China, Korea, Taiwan and Japan 45 were transported to chemical laboratories of IWHR, KKU, National Taiwan University, and CRIEPI, 47 respectively. Each laboratory used its own analytical method. The analytical parameters included pH, conductivity, Na⁺, K⁺, Ca²⁺, Mg²⁺, NH₄⁺, SO₄²⁻, NO₃⁻, 49 and Cl⁻. Since there was no standard solution for 51 chemical analysis of precipitation over East Asia, identical precipitation samples collected in Japan were 53 sent to each laboratory to verify the analytical procedures. For conductivity and pH measurement, samples 55 were analyzed without pretreatment. For other parameters, each sample was filtered through a Milipore

filter with a pore size of 0.45 µm before analysis. The accuracy of chemical analysis was confirmed using 91 two indices, the ratio of total anion to total cation and the ratio of measured to calculated conductivity, 93 based on the method proposed by Miles and Yost (1982). Samples satisfying the condition that 95 deviations of both index values were within +0.2 were identified as 'valid samples'. This checking method was 97 not always implemented for samples obtained in northern China, where relatively high pH values were 99 observed.

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Particle and SO₂ were collected using low-volume 101 samplers with two-stage 47 mm diameter filter packs. The top of the sampler was at approximately 150 cm 103 above the ground level. A Teflon filter was mounted on the upper stage to collect particles, and an alkaline 105 filter on the lower stage to collect SO_2 . The flow rate was set to match the gravitational settlement 107 velocity for a 10 µm spherical particle with a density of 1 g cm^{-3} in the samplers. The alkaline filters were quartz 109 filters impregnated with 6% K₂CO₃ solution for 2h, which were prepared at CRIEPI and mailed to all the 111 stations. The filters were extracted with ultrasonic



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vibration for 1 h in approximately 20 ml of distilled
41 water. The extracted solution was filtered using a Millipore filter and diluted to 50 ml. These filtered
43 samples were analyzed with the same method as precipitation samples.

45 Two 10-day periods (one in each month) were selected for use in the intensive intercomparison excercises.

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49 **3.** Participating models

51 The modeling community was informed of this intercomparison exercise in various ways. The study
53 was announced at various technical meetings, including the 6th Atmospheric Sciences and Application to Air

55 Quality (ASAAQ) Meeting in Beijing (November 1998), and RAINS-Asia Phase II Workshops in Beijing (November 1998) and Bangkok (May 1999). Announcements were also sent to the list of participants at these97meetings and the CRIEPI/IIASA Workshop on Transport of Air Pollutants in Asia (Laxenburg, July 1998).99After the announcement, CDs containing the necessary91input data sets were sent to those responding.101

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Modeling groups that have submitted results to date are listed in Table 4, together with the main character-103 istics of their models and the contributions to the proposed tasks. Models differ substantially from each 105 other, including the computational framework adopted (Eulerian or Lagrangian), the number of vertical layers 107 and transport-diffusion algorithm used. These features are summarized in Tables 5 and 6. Chemical conversion 109 and wet removal parameters adopted by models contributing to the "Blind test" (Task A) are also listed 111 in Table 7. As for the chemical species, all models

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55	tble 5 urther s	odel na	J-SAD	AMS/R	oM-Mo	ΈM	ATCH	ever at a few sites, e.g Possible reasons inclu	ide u
	Ta Fu	М	λ	R	Ю	LS	М	emissions and transp	oort

generally simulate sulfur dioxide and particulate sulfate, 57 linked by a linear transformation and with both removed via dry and wet deposition pathways. A few 59 exceptions exist. The OPU-Model, CRIEPI trajectory and ACDEP-ASIA models treat also sulfate in cloud 61 water as a third transported species. The in-cloud scavenging process is then assumed to take place 63 whether clouds exist or not (clouds are assumed to exist anytime, anywhere), to account for aqueous-phase 65 production of SO₄ in non-precipitating clouds. A corresponding wet removal rate for cloud water sulfate 67 adds a third removal pathway controlled by precipitation. All models used the meteorological data previously 69 described, except MATCH, that used 6-hourly ECMWF data on a lat-lon grid at 1° resolution and 31 hybrid 71 levels in the vertical. 73

4. Results

4.1. Comparison with observations

The initial analysis of the model predictions relative to 79 the observations focused on results from Task A. Sample SO₂ concentration distributions calculated by 81 participating models are shown in Fig. 4 (please note that Model 3 is a receptor oriented model, and did not 83 produce 2D fields). Although absolute values differ from model to model, in all cases the surface SO₂ concentra-85 tions follow closely the emission distribution. The model predicted values are presented along with the observa-87 tions in Figs. 5 and 6 for the 10 day events in January and May. In general the observations fall within the 89 ensemble of the model predictions, with the exception being sulfur-wet deposition at sites like Taichung (which 91 will be discussed later). Furthermore, all the models resolve the observed increase in sulfur deposition (by 93 about one order of magnitude) over east China compared to Japan. The variability of the model 95 predictions is largest over China, and is the smallest for the SO₂ predictions and largest for predictions 97 of sulfur deposition (see Fig. 7 for the bias of model calculated quantities for the month of January). 99 In general the models systematically under-predicted sulfate concentrations and wet sulfur deposition, 101 and over-predicted SO₂ concentrations. This suggests an underestimation of the SO₂ to sulfate conversion 103 rate and will be discussed in further detail in Sections 4.3 and 4.4. 105

The capability of the model ensemble of reproducing the observed values also varies site by site (see Fig. 8). 107 For many locations the ensemble of models has appreciable skill in predicting sulfur deposition. How-109 ever at a few sites, e.g., Oki, all models perform poorly. Possible reasons include uncertainties in regional/local 111 emissions and transport pathways. However, similar

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1 Table 6 Further

Further specifics of Lagrangian models participating to the intercomparison

	Layers	Trajectories type	Horizontal diffusion	Vertical diffusion
CRIEPI Traj. Model	1 layer: BL (1000 m)	backward (forward Euler procedure) on isobaric winds at 850 hPa or at the bottom σ_z , when surface pressure of the travelling air mass is below 850 hPa	Gaussian	uniform within the BL
ACDEP ASIA	1 layer: BL (variable height)	backward (Petterssen's procedure) on isosigma winds at $\sigma_r = 1000 \text{ m}$	Gaussian, with $\sigma_h = 0$. (m) $d =$ travel distance	1 <i>d</i> uniform within the BL (m)
ATMOS-2	2 or 3 layers: BL (variable height), free troposphere (6000 m), nighttime surface layer	forward (iterative forward Euler procedure) on layer- averaged winds	Gaussian, with $\sigma_h = 0$. (m) $t =$ travel time (s)	5t uniform within each layer mass exchanges among layers at sunrise and sunse
	(300 m)			
Table 7 Models-specif	ic parameters for Task A ^a			
Model code	$SO_2 \rightarrow SO_4$ conversion rate	(s^{-1}) SO ₂ wet re	emoval rate (s ⁻¹)	SO_4 wet removal rate (s ⁻¹)
1	$k_c = \bar{k}[1 - 0.4\cos(2\pi H/24)]$] $4 \times 10^{-5} P$		$10 \times 10^{-5} P$
	$\bar{k} = k_{EQ} + \frac{ \lambda }{90}(k_{\text{POLE}} - k_{EQ})$			
	$k_{EQ} = 4.0 \times 10^{-6}$			
	$k_{\rm POLE} = 1.3 \times 10^{-6} + 1.1 \times$	$10^{-6} \sin \gamma$		
	$\int \frac{2\pi(J+91)}{\lambda}, \lambda < 0$			
	$\gamma = \begin{cases} \frac{365}{2\pi(J-91)} & \lambda \ge 0 \end{cases}$			
	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$	av. 1 = lotitude		
	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$	ay, $\lambda = $ latitude		
2	$\gamma = \begin{cases} \frac{365}{2\pi(J-91)}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d}$ (Tarrasón, 1995) $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6}$	ay, $\lambda = $ latitude $2 \times 10^{-5}P$		$2 imes 10^{-4} P^{0.7}$
2	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$ $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6} \\ \max\left(0; \frac{\sin(\theta(\lambda, \omega))}{\sin(\theta(50^\circ \text{N}, 25^\circ \text{N}, $	ay, $\lambda = $ latitude $2 \times 10^{-5} P$ $\frac{(0)}{(1 \text{ June})}$		$2 \times 10^{-4} P^{0.7}$ (Maeda, 1998)
2	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$ $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6} \\ \max\left(0; \frac{\sin(\theta(\lambda, \lambda))}{\sin(\theta(50^\circ \text{N}, 22))}\right)$ $\theta = \text{height of the sun at mix}$ $\lambda = \text{latitude, } J = \text{Julian day}$	ay, $\lambda = $ latitude $2 \times 10^{-5} P$ $\frac{())}{(1 - 5)}$ lday,		$2 \times 10^{-4} P^{0.7}$ (Maeda, 1998)
2	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$ $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6} \\ \max\left(0; \frac{\sin(\theta(\lambda, z))}{\sin(\theta(50^\circ\text{N}, 2z))}\right)$ $\theta = \text{height of the sun at mic} \\ \lambda = \text{latitude, } J = \text{Julian day} \\ (\text{Christensen 1995}) \end{cases}$	ay, $\lambda = $ latitude $2 \times 10^{-5} P$ $\frac{(2)}{(3 \text{ June}))$ Iday,		2 × 10 ⁻⁴ P ^{0.7} (Maeda, 1998)
2 3 and 4	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$ $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6} \\ \max\left(0; \frac{\sin(\theta(\lambda_c), \lambda_c)}{\sin(\theta(50^\circ \text{N}, 22))}\right)$ $\theta = \text{height of the sun at mic} \\ \lambda = \text{latitude, } J = \text{Julian day} \\ (\text{Christensen 1995}) \end{cases}$ $2.8 \times 10^{-6} \text{ in January} \\ 8.5 \times 10^{-5} \text{ in May} \\ 5.6 \times 10^{-6} \text{ rainout} \end{cases}$	ay, $\lambda = $ latitude $2 \times 10^{-5} P$ ()) λ June))) lday, 2.8×10^{-5}	P below-cloud	$2 \times 10^{-4} P^{0.7}$ (Maeda, 1998) $2.8 \times 10^{-5} P$ below-cloud $2.8 \times 10^{-5} P$ for cloud SO ₄
2 3 and 4	$\gamma = \begin{cases} 365\\ \frac{2\pi(J-91)}{365}, & \lambda \ge 0 \end{cases}$ $H = \text{local time, } J = \text{Julian d} \\ (\text{Tarrasón, 1995}) \end{cases}$ $k_c = 0.3 \times 10^{-6} + 4.7 \times 10^{-6} \\ \max\left(0; \frac{\sin(\theta(\lambda, \cdot, \sin(\theta(50^\circ\text{N}, 2: \theta - 6))))) \right)$ $\theta = \text{height of the sun at mix} \\ \lambda = \text{latitude, } J = \text{Julian day} \\ (\text{Christensen 1995}) \\ 2.8 \times 10^{-6} \text{ in January} \\ 8.5 \times 10^{-5} \text{ in May} \\ 5.6 \times 10^{-6} \text{ rainout} \\ 7.4 \times 10^{-6} \text{ in summer} \end{cases}$	ay, $\lambda = $ latitude $2 \times 10^{-5} P$ $\frac{f(y)}{3 \text{ June}(y)}$ lday, 2.8×10^{-5} 3.9×10^{-5}	P below-cloud $P^{0.12}$ in summer	$2 \times 10^{-4} P^{0.7}$ (Maeda, 1998) $2.8 \times 10^{-5} P$ below-cloud $2.8 \times 10^{-5} P$ for cloud SO ₄ $1.1 \times 10^{-4} P^{0.06}$ in summer

^aModels 5, 7, and 8 used the same parameters as presented in Table 2.

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performance problems are not seen at nearby sites,
suggesting that reasons may be related to horizontal resolution issues, and these were not addressed in this
phase of the study.

There are many factors that influence the model results. One very important factor is the direct relationship between sulfate wet deposition and precipitation 73 rate and amount. The analysis of model predicted precipitation against observed amounts (see Fig. 3) 55 showed that the largest underpredictions by the models occurred for cases where the modeled precipitation was underpredicted, and these occurred mostly over the island stations. For this reason measured precipitation 103 amounts were used to adjust modeled wet depositions before the comparison with observations (and through 105 this, model predicted values were found to follow more closely with observations). Furthermore, model perfor-107 mance was found to increase when the comparisons were made on longer time periods (e.g. on a monthly 109 basis; not shown), as a result of the fact that the monthly time frame captured and summed the influence of more 111 precipitation events.

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45 4.2. Transport and diffusion

47 Each model differs in important ways including the numerical algorithm used, model framework, and the
49 vertical resolution (see Tables 5 and 6). Each of these features impacts the resulting calculated spatial and
51 temporal distributions of sulfur concentrations and deposition. To illustrate this point calculated daily mean
53 SO₂ concentrations for 30 May are presented in Figs. 9a and b for two regions: South Korea and southern
55 China (see Fig. 4 for the region locations). The daily mean near-surface concentrations and the boundary

layer (0–1 km) averaged distributions computed by 101 all models with fixed parameters (Task B) are shown. All models show the same general features, with the 103 high emission areas clearly located within the subdomains. However the models do differ in terms of 105 the magnitudes of the peak values, and the structure of the horizontal and vertical distributions. For 107 example, in southern Korea, Model 1 shows much stronger horizontal and weaker vertical gradients than 109 Model 5. Models 1 and 8, followed by 2 and 7, also present in general the highest surface maximum 111 (see also Fig. 10).



The differences between model predictions are more clearly depicted in Fig. 11, where the ratio of predicted near-surface to boundary-layer concentrations at various locations are presented. Again the general behavior of all the models is similar, with all models resolving strong vertical gradients in the high source areas such as Komae and Nanjing, and small gradients over island locations with no appreciable anthropogenic SO₂ emissions (e.g., Oki, Amami, Miyako). The only point with

boundary layer concentrations higher than surface is at Hachijo, a location heavily impacted by an elevated 105 strong volcanic emission. However the magnitudes of this ratio varies by as much as a factor of 5. Consistent 107 with what was already pointed out for concentrations integrals, Models 2 and 7 generally have higher surface/ 109 boundary layer concentration ratios, meaning that they exhibit sharper vertical gradients, while the opposite 111 holds for Model 1, which generally shows a relatively

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Fig. 9. (a) SO₂ daily surface (top row) and boundary layer averaged concentrations (bottom row) on 30 May over southern Korea
 (contours at 1, 2, 5, 10, 20, 30, 40... ppb). Note that Model 3 is receptor oriented model and did not produce 2D fields, Models 4 and 8 are 1-layer models so boundary layer averaged plots (not shown) are the same as the surface layer plots. (b) Same as (a), but for Central
 China (contours at 1, 2, 5, 10, 15, 20, 30, 40... ppb).



53 modest vertical gradient in the lower part of the troposphere.

55 This point is amplified further in Fig. 12, where the masses of SO_2 as sulfur in the lowest 1 km for each

of the target regions are presented. Again all models 109 show the same general features in terms of which region has the highest mass loadings (i.e., Central China), but 111 the magnitude of the mass can vary by a factor of 5.

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Fig. 11. SO₂ daily concentrations at monitoring stations (30 May): ratio between surface values and the average over the first 1000 m.



Fig. 12. Integrals in the first 1000 m of SO₂ daily concentrations over the four target regions (30 May).

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41 Similar findings hold for January results as well (not shown). Such differences are less dependent on
43 model framework (Lagrangian vs. Eulerian), than on the number of vertical model layers, and the
45 numerical algorithms (i.e., diffusive characteristics) and assumptions (e.g., whether horizontal diffusion was
47 included explicitly or not) used for horizontal and vertical transport.

4.3. Chemical conversion

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The conversion rate of SO₂ to sulfate is a key
53 parameter in determining the partitioning of sulfur between SO₂ and sulfate, and in turn, in determining the
55 pathways for deposition (in what form and by what process—wet/dry), and the overall lifetime of sulfur in

the atmosphere. One metric that reflects overall conver-77 sion rates in the context of all removal processes is the fraction of sulfur in ambient air present in the form of 79 sulfate. A plot of observed and modeled predicted fraction of sulfur as sulfate in ambient air ([SO₄]/ 81 $([SO_2] + [SO_4])$, both species as S) is presented in Figs. 13a and b. Shown are the values calculated by 83 models using their own parameters (Task A). The low value at Amami in May is the result of measurement 85 error with ambient values at/or below the instrument detection limit. These results show that the fraction of 87 sulfur as sulfate is in the range of 0.3-0.4 over Japan, and increases to 0.6-0.8 over China and Korea. The 89 model predicted values are generally consistent with the observations and exhibit higher (or lower) ratios in both 91 months over the same regions.

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One interesting feature is that the observations in 93 general do not show a dramatic increase in the fraction as sulfate in May relative to that in January. Based 95 simply on chemical conversion rates, one would expect a marked difference (with the fraction as sulfate in May 97 higher than that in January) as the gas phase conversion rates are about three times higher in May than January. 99 The fact that a large increase is not apparent demonstrates the tight coupling between chemical conversion 101 and transport and removal processes. If the transport and removal characteristics were identical, then based 103 on chemical conversion considerations only, May values of sulfate concentration and sulfate fraction should be 105 higher than those in January, while SO₂ concentrations in January should be higher than those in May. As also 107 shown in Figs. 13c and d there is no clear trend in SO₂ and sulfate levels between May and January. This 109 suggests that the transport and removal characteristics of the events may be different and that they play an 111 unimportant role. The monthly mean wind speeds and

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Fig. 13. Fraction of SO₄ in air concentrations: (a) 11–20 January; (b) 21–30 May. January/May ratio calculated for 10 days average
 concentrations: (c) SO₂; (d) SO₄ (Task A).

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Fig. 14. Total sulfur depositions for both months (January and May together) and contributions due to each process over selected target regions. Upper panels: results from Task B. Lower panels: Task A vs. Task B.

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precipitation (Fig. 2) show that the transport speeds and
general precipitation patterns are quite similar for these
two periods over most of the domain. However, in the
case of precipitation the number of events was greater in
January, while the intensities were higher in May.
Furthermore the precipitation amounts over the high
sulfur source regions in eastern China were higher in
May. The combinations of these factors control the net
sulfate to sulfur ratio.

A A Sulfur

4.4. Sulfur deposition

The model calculated sulfur deposition was further 47 studied by comparing the sulfur deposition quantities and the contributions due to each process (i.e., dry/wet 49 as SO₂ and sulfate) at the four target regions. Examples of the results compiled for the 2-month period (January and May together) are shown in the upper panels of 51 Fig. 14, for South Korea and Central China. While variations in predicted sulfur deposition range from 53 $\sim 50\%$ in South Korea to 400% in Central China, the 55 predictions are very similar in terms of the relative importance of the various pathways. Dry deposition as



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Fig. 15. Wet to total sulfur ratio in depositions for both months (January and May together) over the target regions (Task A). 107

SO₂ and wet deposition as sulfate are the two dominant terms. In fact the models are very robust in predicting the wet to total sulfur deposition ratio. This is shown in Fig. 15, where all the models predict that the fraction

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- 1 wet removed over Japan to be between 0.6 and 0.7 and to drop to between 0.35 and 0.5 over eastern China. The
- 3 predictions of absolute quantities can vary greatly as shown over Central China. Here the differences also
- 5 show up in the various terms, with Model 2 showing a large wet sulfate contribution (this model also has high
- 7 sulfate gas phase concentrations). The increased difference among models over Central China may be
- 9 attributed to the combination of high emission fluxes and low ventilation (see Fig. 2).
- 11 It is interesting to assess whether these differences are due to model structure or to differences in the chemical
- 13 conversion and wet removal parameters. A comparison of model results for Tasks A and B allow a direct
- 15 evaluation of how much of these differences can be attributed to the chemical and removal parameters used
- in the model calculations. These results are shown in the lower panels of Fig. 14 for the models that performedboth tasks, using different parameters while the para-
- meters used are summarized in Table 7. Relative to the fixed parameter case the SO_2 to sulfate conversion rates

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varied from 0% to -100%, while the wet removal rates varied by factors of -10 to +2. Quite surprisingly, the results from Tasks A and B are very consistent, indicating that the choice of parameters within this range of values has a remarkably *small* impact on the results, and that the biggest influence is in terms of model structure. 63

Another key issue in analyzing model performance in comparison with observations is the issue of horizontal 65 resolution. Long-range transport models calculate concentrations and deposition amounts relative to a 67 horizontal grid, while a measurement site obtains values at a fixed point. This fact gives rise to the classical 69 problem of how representative is a point measurement to the value within a model grid cell, and how to take 71 into account this sub-grid scale variability in model analysis. This point is illustrated in Fig. 16, where 73 the measurements sites of Kanghwa and Tsushima are located on SO₂ emissions with resolutions of 1° and 30 s. 75 Kwanghwa and Tsushima when placed within the 1° emissions grid represent sites located within the highest 77



Fig. 16. Monitoring station location and emission distribution at Kanghwa (left) and Tsushima (right); spatial resolution is 1° on 111 upper row and 30 s on bottom row.

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and 4th lowest emissions, respectively. The models 1 consistently over-predicted the concentrations (both sulfur dioxide and sulfate) at these two sites 3 (e.g., see May values), but do better with sulfur 5 wet deposition. These features can be understood from Fig. 16, where Tsushima is shown to be isolated 7 from local sources for all flow directions, and Kanghwa is isolated from local sources for all flow directions 9 except easterly. Sulfate wet deposition is a more integrated quantity than surface concentrations. 11 and is more heavily influenced by long-range transport, and more reflective of larger scale features. Engardt (2001) in the course of this study explored 13

various ways of interpolating model results to estimate
values at specific monitoring site locations. Bi-linear
interpolation from surrounding nearest neighbor grid
points to the observation site location was found to
improve model performance compared to simply
reporting the value in the grid where the measurement
site is located.

4.5. Source-receptor relationships

To evaluate how model formulation and choice of
parameters affect the source-receptor relationships,
MICS participants were asked to provide source-
receptor relationships at a subset of locations for the
set of source regions, shown in Fig. 1. The results are
presented in Figs. 17 and 18 for January and May,
respectively.63

The models show in many instances a remarkable consistency in terms of identifying the major source 67 contributor. For example at Komae in January all models show the contribution from Japan sources to be 69 the greatest single contributor with contributions ranging from 40% to 60%. In both months the contribution 71 of Central eastern China at Beijing is above 60%, and in Naniing above 80%. Nearby sources are also the 73 greatest contributors to depositions at Yangyang and Taichung, especially in May, with contributions ranging, 75 respectively, from 40% to 60% and from 60% to 90%.



55 Fig. 19. (a) 5 day iso-sigma back-trajectories arriving in Taichung. Shown are the frequency distribution of all 3-h end points over the 111 entire month of trajectories originated at the receptor. (b) As in (a), but for Fukue.

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35 Noticeable differences do arise; for example the variation in the contribution due to Central eastern 37 China to the receptor at Taichung, and the great structure in contributions to the deposition at Fukue 39 in May. Based on the discussions in Section 4.4, these differences are likely due to transport and vertical 41 resolution issues and not choices of the chemical and deposition parameters. To test this hypothesis statistics 43 of back trajectories were calculated for each receptor and the impact region identified. Trajectories were 45 initiated every 3h and followed for 5 days. The probability distribution was constructed using all 3-h 47 end-point locations for the entire month. Thus the highest probabilities are near the receptor. Moving away 49 from the receptor reflects going farther back in time. Figs. 19a and b show examples of trajectory statistics at different levels for two of the receptors. The plots for 51 Taichung help explain why the models generally agree in 53 terms of source-receptor relationships in May but differ

in January. The effect of the WM on transport (and thus source-receptor relationships) is clearly depicted, with

55 source-receptor relationships) is clearly depicted, with transport in the lower 1 km coming to Taichung either

from Central eastern China or from out over the ocean, 91 while the transport above this level originates from the southern China source regions. Variations in the relative 93 contributions of these two sources is the result of transport layers, which in turn depends on model 95 resolution as well as the use of forward or backward trajectories. For May the transport patterns are much 97 less dependent on height. For Fukue the large number of contributing sources and the relative differences are due 99 largely to the fact that this receptor is more centrally located and surrounded by a large number of source 101 regions, and that the trajectories originate from generally wider areas. 103

4.6. Sensitivity studies 105

The discussions in this paper have focused on the analysis of results from the participating models for the specific tasks assigned. However, many of the participating modeling groups performed additional studies to better understand how their models performed under the conditions specified. These investigations took the form

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- 1 of additional systematic studies exploring model sensitivity to the number of vertical layers and the height of
- 3 the near surface layers, various methods of estimating horizontal and vertical mixing, and different chemical
- 5 conversion and wet removal rates (Lee and Kim, 2000; Hayami and Ichikawa, 2001; Engardt, 2001). These
- 7 results indicate that the major factors affecting model performance are the emissions, the driving meteorolo-
- 9 gical conditions, and model formulation factors related to vertical diffusion and mixing height. Vertical resolu-
- 11 tion was identified as an important factor in diagnosing the depth of the daytime mixed-layer. The sensitivity of
- 13 the models to these aspects of formulation and parameterization are largest in high emission regions and
- 15 over land (where diurnal mixing variations are largest), and help explain why the MICS models showed the
- 17 largest differences over Central China.

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5. Concluding remarks

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An intercomparison study of long-range transport 23 models in East Asia has been initiated. This exercise has promoted collaborative work in the region and yielded 25 some interesting and important results. For example, the difference in results between Task A and B parameters 27 were found not to differ very substantially in terms of

- concentrations and depositions, although the chemicalconversion and scavenging rates differ by at least a factor of 2. At this stage we judge the uncertainty in
- 31 emission inventory, and secondarily the driving meteorology, to be much larger than the uncertainty in the
- model parameters. As the combined results from the different models and tasks have shown, it is in fact
 difficult to adjust the parameters of these models to
- cover the broad spectrum of conditions occurring withinthe domain in different seasons. This problem is also strengthened by the objective difficulties in collecting
- 39 coherent sets of long-term measurements over the whole area.
- Finally, it needs to be emphasized that the main aim of the study was not to rank individual model
 performances according to their agreement against measurements. Although this is very important, in the
 current phase of the study attention was rather focused on estimating the uncertainties of predictions made by
- 47 the set of participating models, and understanding the reasons for the resulting similarities and discrepancies.
- 49 The results from this study provide an idea about the level of confidence that can be expected from current
- 51 applications of long-range transport models to East Asia.
- 53 Clearly, more work and analysis is needed. The intercomparison study is on going and we invite and
- 55 encourage others to participate (http://www.cgrer.uiowa.edu/people/gcalori/model_intercomp.htm). There is

a plan to expand the focus to include annual deposition 57 and nitrate deposition. 59

6. Uncited References

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1997.	65

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